# Energy and potential enstrophy flux constraints in quasi-geostrophic models

#### **Eleftherios Gkioulekas**

University of Texas Rio Grande Valley

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### Publications

- ▶ K.K. Tung and W.W. Orlando (2003a), J. Atmos. Sci. 60,, 824-835.
- ► K.K. Tung and W.W. Orlando (2003b), Discrete Contin. Dyn. Syst. Ser. B, 3, 145-162.
- ▶ K.K. Tung (2004), J. Atmos. Sci., 61, 943-948.
- ► E. Gkioulekas and K.K. Tung (2005), *Discr. Cont. Dyn. Sys. B* 5, 79-102
- ▶ E. Gkioulekas and K.K. Tung (2005), *Discr. Cont. Dyn. Sys. B* 5, 103-124.
- ► E. Gkioulekas and K.K. Tung (2007): J. Fluid Mech., 576, 173-189.
- ► E. Gkioulekas and K.K. Tung (2007): Discrete Contin. Dyn. Syst. Ser. B, 7, 293-314
- ► E. Gkioulekas (2012): J. Fluid Mech. 694, 493-523
- ▶ E. Gkioulekas (2014): Physica D 284, 27-41

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## Outline

- ► Flux inequality for 2D turbulence
- ► Flux inequality for multi-layer symmetric QG model
  - Formulation
  - Dissipation rate spectra
  - symmetric streamfunction dissipation
- ► Flux inequality for two-layer QG model
  - asymmetric streamfunction dissipation
  - differential diffusion
  - extrapolated Ekman term

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#### 2D Navier-Stokes equations

► In 2D turbulence, the scalar vorticity  $\zeta(x, y, t)$  is governed by

$$\frac{\partial \zeta}{\partial t} + J(\psi, \zeta) = d + f,$$

where  $\psi(x, y, t)$  is the streamfunction, and  $\zeta(x, y, t) = -\nabla^2 \psi(x, y, t)$ , and

$$d = -[\nu(-\Delta)^{\kappa} + \nu_1(-\Delta)^{-m}]\zeta$$

 The Jacobian term J(ψ, ζ) describes the advection of ζ by ψ, and is defined as

$$J(a,b) = \frac{\partial a}{\partial x} \frac{\partial b}{\partial y} - \frac{\partial b}{\partial x} \frac{\partial a}{\partial y}$$

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## Energy and enstrophy spectrum. I

Two conserved quadratic invariants: energy E and enstrophy G defined as

$$E(t) = -\frac{1}{2} \int \psi(x,y,t) \zeta(x,y,t) \, \mathrm{d}x \mathrm{d}y \quad G(t) = \frac{1}{2} \int \zeta^2(x,y,t) \, \mathrm{d}x \mathrm{d}y.$$

Let a<sup><k</sup>(x) be the field obtained from a(x) by setting to zero, in Fourier space, the components corresponding to wavenumbers with norm greater than k:

$$\begin{aligned} a^{$$

Filtered inner product:

$$\langle a,b\rangle_k = \frac{\mathrm{d}}{\mathrm{d}k} \int_{\mathbb{R}^2} \mathrm{d}\mathbf{x} \; a^{< k}(\mathbf{x}) b^{< k}(\mathbf{x})$$

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## Energy and enstrophy spectrum. II

- Energy spectrum:  $E(k) = -\langle \psi, \zeta \rangle_k$
- Enstrophy spectrum  $G(k) = \langle \zeta, \zeta \rangle_k$
- Consider the conservation laws for E(k) and G(k):

$$\frac{\partial E(k)}{\partial t} + \frac{\partial \Pi_E(k)}{\partial k} = -D_E(k) + F_E(k)$$
$$\frac{\partial G(k)}{\partial t} + \frac{\partial \Pi_G(k)}{\partial k} = -D_G(k) + F_G(k)$$

► In two-dimensional turbulence, the energy flux  $\Pi_E(k)$  and the enstrophy flux  $\Pi_G(k)$  are constrained by

$$k^2 \Pi_E(k) - \Pi_G(k) \le 0$$

#### for all *k* not in the forcing range.

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## Energy and enstrophy spectrum. III

Assuming a forced-dissipative configuration at steady state,

$$\begin{split} \Pi_E(k) &= \int_k^{+\infty} D_E(q) \mathrm{d} q, \\ \Pi_G(k) &= \int_k^{+\infty} D_G(q) \mathrm{d} q, \end{split}$$

and it follows that:

$$k^{2}\Pi_{E}(k) - \Pi_{G}(k) = \int_{k}^{+\infty} [k^{2}D_{E}(q) - D_{G}(q)] dq = \int_{k}^{+\infty} \Delta(k,q) dq.$$

► For the case of two-dimensional Navier-Stokes turbulence,  $D_G(k) = k^2 D_E(k)$ , therefore

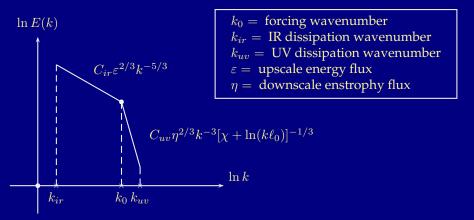
$$\Delta(k,q) = k^2 D_E(q) - D_G(q) = (k^2 - q^2) D_E(q) \le 0$$

so we get 
$$k^2 \Pi_E(k) - \Pi_G(k) < 0$$
.

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## KLB theory.



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### **Cascade Directions**

- Proofs that energy goes mostly upscale in 2D turbulence:
  - ► Fjørtøft (1953): Famous wrong argument.
  - Merilee and Warn (1975): Noticed error in Fjørtøft
  - Eyink (1996): Correct argument, but assumes inertial ranges.
  - ► Gkioulekas and Tung (2007): Flux inequality for 2D Navier-Stokes
- Linear Cascade Superposition hypothesis:
  - ► E. Gkioulekas and K.K. Tung (2005), *Discr. Cont. Dyn. Sys. B* 5, 79-102
  - E. Gkioulekas and K.K. Tung (2005), Discr. Cont. Dyn. Sys. B 5, 103-124.

Under coexisting downscale cascades of energy and enstrophy:

$$E(k) \approx C_1 \varepsilon_{uv}^{2/3} k^{-5/3} + C_2 \eta_{uv}^{2/3} k^{-3}$$

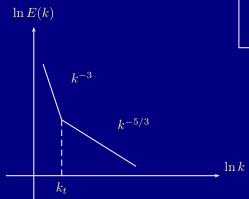
with  $\eta_{uv}$  the downscale enstrophy flux and  $\varepsilon_{uv}$  the downscale energy flux.

• Transition wavenumber:  $k_t = \sqrt{\eta/\varepsilon}$ .

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#### Nastrom-Gage spectrum schematic

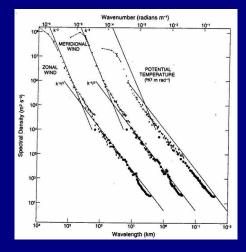


 $\begin{array}{l} k^{-3} \rightarrow 3000 \mathrm{km} - 800 \mathrm{km} \\ k^{-5/3} \rightarrow 600 \mathrm{km} - \ll 1 \mathrm{km} \\ k_t \approx 700 \mathrm{km} \approx k_R \end{array}$ 

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#### Nastrom-Gage spectrum

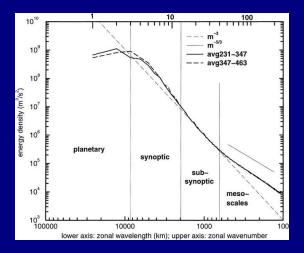


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### Tung and Orlando spectrum



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#### Generalized multi-layer model. I.

Consider the generalized form of an n-layer model:

$$\frac{\partial q_{\alpha}}{\partial t} + J(\psi_{\alpha}, q_{\alpha}) = d_{\alpha} + f_{\alpha}$$
$$d_{\alpha} = \sum_{\beta} \mathcal{D}_{\alpha\beta}\psi_{\beta}$$
$$\dot{t}_{\alpha}(\mathbf{k}, t) = \sum_{\beta} L_{\alpha\beta}(\|\mathbf{k}\|)\hat{\psi}_{\beta}(\mathbf{k}, t)$$

- We consider two types of forms for the dissipation term  $d_{\alpha}$ :
  - Let  $D_{\alpha}(k)$  be the spectrum of the operator  $\mathscr{D}_{\alpha}$
  - Streamfunction dissipation:
    - $d_{\alpha} = +\mathscr{D}_{\alpha}\psi_{\alpha} \Longrightarrow D_{\alpha\beta}(k) = \delta_{\alpha\beta}D_{\beta}(k)$
  - Symmetric streamfunction dissipation: (all layers have the same operator)

$$\hat{d_{\alpha}} = +\mathscr{D}\psi_{\alpha} \Longrightarrow D_{\alpha\beta}(k) = \delta_{\alpha\beta}D(k)$$

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#### Generalized multi-layer model. II

► The energy spectrum *E*(*k*) and the potential enstrophy spectrum *G*(*k*) are given by:

$$E(k) = -\sum_{\alpha} \langle \psi_{\alpha}, q_{\alpha} \rangle_{k} = -\sum_{\alpha\beta} L_{\alpha\beta}(k) C_{\alpha\beta}(k)$$
$$G(k) = \sum_{\alpha} \langle q_{\alpha}, q_{\alpha} \rangle_{k} = \sum_{\alpha\beta\gamma} L_{\alpha\beta}(k) L_{\alpha\gamma}(k) C_{\beta\gamma}(k)$$

with  $C_{\alpha\beta}(k) = \langle \psi_{\alpha}, \psi_{\beta} \rangle_{k}$ .

► The energy dissipation rate spectrum D<sub>E</sub>(k) and the layer-by-layer potential enstrophy dissipation rate spectra D<sub>G<sub>α</sub></sub>(k) are given by

$$\begin{split} D_E(k) &= 2\sum_{\alpha\beta} D_{\alpha\beta}(k) C_{\alpha\beta}(k), \\ D_{G_{\alpha}}(k) &= -2\sum_{\beta\gamma} L_{\alpha\beta}(k) D_{\alpha\gamma}(k) C_{\beta\gamma}(k). \end{split}$$

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## Multi-layer QG model. I.

 In a multi-layer quasigeostrophic model, the relation between q<sub>α</sub> and ψ<sub>α</sub> reads

$$\begin{split} q_1 &= \nabla^2 \psi_1 + \mu_1 k_R^2 (\psi_2 - \psi_1) \\ q_\alpha &= \nabla^2 \psi_\alpha - \lambda_\alpha k_R^2 (\psi_\alpha - \psi_{\alpha-1}) + \mu_\alpha k_R^2 (\psi_{\alpha+1} - \psi_\alpha), \text{ for } 1 < \alpha < n \\ q_n &= \nabla^2 \psi_n - \lambda_n k_R^2 (\psi_n - \psi_{n-1}) \end{split}$$

• Here  $\lambda_{\alpha}, \mu_{\alpha}$  are the non-dimensional Froude numbers given by

$$\lambda_{\alpha} = \frac{h_1}{h_{\alpha}} \frac{\rho_2 - \rho_1}{\rho_{\alpha} - \rho_{\alpha-1}}, \text{ for } 1 < \alpha \le n$$
$$\mu_{\alpha} = \frac{h_1}{h_{\alpha}} \frac{\rho_2 - \rho_1}{\rho_{\alpha+1} - \rho_{\alpha}}, \text{ for } 1 \le \alpha < n$$

with  $h_1, h_2, \ldots, h_n$ , the thickness of layers from top to bottom, in pressure coordinates.

▶ For  $h_1 = h_2 = \ldots = h_n$ , we note that  $\lambda_{\alpha+1} = \mu_\alpha$  for all  $1 \le \alpha < n$ .

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## Multi-layer QG model. II.

• The corresponding matrix  $L_{\alpha\beta}(k)$  is given by:

$$L_{\alpha\alpha}(k) = \begin{cases} -k^2 - \mu_1 k_R^2, & \text{if } \alpha = 1\\ -k^2 - (\lambda_\alpha + \mu_\alpha) k_R^2, & \text{if } 1 < \alpha < n\\ -k^2 - \lambda_n k_R^2, & \text{if } \alpha = n \end{cases}$$
$$L_{\alpha,\alpha+1}(k) = \mu_\alpha k_R^2, \text{ for } 1 \le \alpha < n$$
$$L_{\alpha-1,\alpha}(k) = \lambda_\alpha k_R^2, \text{ for } 1 < \alpha \le n$$

► We define:

$$\gamma_{\alpha}(k,q) = k^2 + \sum_{\beta} L_{\alpha\beta}(q) = k^2 - q^2 < 0, \text{ for } k < q$$

► We consider the case where  $h_{\alpha} = h$  for all layers. Then,  $L_{\alpha\beta}(k)$  is symmetric, and our theoretical framework becomes applicable.

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## Multi-layer QG model. III.

- ► Let  $U_{\alpha}(k) = \langle \psi_{\alpha}, \psi_{\alpha} \rangle_{k} \ge 0$  for  $1 \le \alpha \le n$ . (streamfunction spectrum)
- Recall that:

$$k^{2}\Pi_{E}(k) - \Pi_{G}(k) = \int_{k}^{+\infty} [k^{2}D_{E}(q) - D_{G}(q)] dq = \int_{k}^{+\infty} \Delta(k,q) dq.$$

▶ PROPOSITION 1: In a generalized *n*-layer model, under symmetric streamfunction dissipation  $d_{\alpha} = +\mathscr{D}\psi_{\alpha}$  with spectrum D(k), we assume that  $L_{\alpha\beta}(q) \ge 0$  when  $\alpha \ne \beta$ , and  $L_{\alpha\beta}(q) = L_{\beta\alpha}(q)$ , and  $\gamma_{\alpha}(k,q) \le 0$  when k < q for all  $\alpha$ . It follows that:

$$\Delta(k,q) \le D(q) \sum_{\alpha} \gamma_{\alpha}(k,q) U_{\alpha}(q) \le 0$$

- It follows that under symmetric streamfunction, the flux inequality is satisfied.
- ► The case d<sub>α</sub> = Dq<sub>α</sub> presents unexpected challenges, and may violate the flux inequality in models with more than 2 layers.

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### The two-layer model. I

 The governing equations for the two-layer quasi-geostrophic model are

$$\begin{aligned} \frac{\partial \zeta_1}{\partial t} + J(\psi_1, \zeta_1 + f) &= -\frac{2f}{h}\omega + d_1 \\ \frac{\partial \zeta_2}{\partial t} + J(\psi_2, \zeta_2 + f) &= +\frac{2f}{h}\omega + d_2 \\ \frac{\partial T}{\partial t} + \frac{1}{2}[J(\psi_1, T) + J(\psi_2, T)] &= -\frac{N^2}{f}\omega + Q_0 \end{aligned}$$

where  $\zeta_1 = \nabla^2 \psi_1$ ;  $\zeta_2 = \nabla^2 \psi_2$ ;  $T = (2/h)(\psi_1 - \psi_2)$ . *f* is the Coriolis term; *N* the Brunt-Väisälä frequency;  $Q_0$  is the thermal forcing on the temperature equation;  $d_1$ ,  $d_2$  the dissipation terms.

- ► The three equations are situated in three layers:
  - $\psi_1$ : At 0.25Atm, top streamfunction layer
  - ► *T*: At 0.5Atm, temperature layer.
  - $\psi_2$ : At 0.75Atm, bottom streamfunction layer

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#### The two-layer model. II

The potential vorticity is defined as

$$q_1 = \nabla^2 \psi_1 + f + \frac{k_R^2}{2} (\psi_2 - \psi_1)$$
$$q_2 = \nabla^2 \psi_2 + f - \frac{k_R^2}{2} (\psi_2 - \psi_1)$$

with  $k_R \equiv 2\sqrt{2}f/(hN)$  and it satisfies

$$\frac{\partial q_1}{\partial t} + J(\psi_1, q_1) = f_1 + d_1$$
$$\frac{\partial q_2}{\partial t} + J(\psi_2, q_2) = f_2 + d_2 + e_2$$

with  $f_1 = (1/4)k_R^2 h Q_0$  and  $f_2 = -(1/4)k_R^2 h Q_0$ .

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#### The two-layer model. III

We use the following asymmetric dissipation configuration:

$$d_1 = \nu(-1)^{p+1} \nabla^{2p+2} \psi_1, \tag{1}$$

$$d_2 = (\nu + \Delta \nu)(-1)^{p+1} \nabla^{2p+2} \psi_2 - \nu_E \nabla^2 \psi_s.$$
 (2)

- Differential hyperdiffusion:  $\Delta \nu > 0$ .
- The Ekman term is given in terms of the streamfunction ψ<sub>s</sub> at the surface layer (p<sub>s</sub> = 1Atm) which is linearly extrapolated from ψ<sub>1</sub> and ψ<sub>2</sub> and it is given by ψ<sub>s</sub> = λψ<sub>2</sub> + μλψ<sub>1</sub>, with λ and μ given by

$$\lambda = \frac{p_s - p_1}{p_2 - p_1} \text{ and } \mu = \frac{p_2 - p_s}{p_s - p_1}.$$
(3)

$$\blacktriangleright \ 0 < p_1 < p_2 < p_s \Longrightarrow -1 < \mu < 0$$

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#### The two-layer model. IV

 The dissipation term configuration corresponds to setting the generalized dissipation operator spectrum D<sub>αβ</sub>(k) equal to

$$D(k) = \begin{bmatrix} D_1(k) & 0\\ \mu d(k) & D_2(k) + d(k) \end{bmatrix},$$
 (4)

with  $D_1(q)$ ,  $D_2(q)$ , and d(q) given by

$$D_1(k) = \nu k^{2p+2}$$
 and  $D_2(k) = (\nu + \Delta \nu) k^{2p+2}$  and  $d(k) = \lambda \nu_E k^2$ .  
(5)

• The nonlinearity corresponds to an operator  $\mathcal{L}_{\alpha\beta}$  with spectrum  $L_{\alpha\beta}(k)$  given by

$$L(k) = -\begin{bmatrix} a(k) & b(k) \\ b(k) & a(k) \end{bmatrix},$$
(6)

with a(k) and b(k) given by  $a(k) = k^2 + k_R^2/2$  and  $b(k) = -k_R^2$ .

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#### The two-layer model. V

► PROPOSITION 2:Assume streamfunction dissipation with both differential small-scale dissipation and extrapolated Ekman dissipation with  $-1 < \mu < 0$ . Assume also that  $k^2 - a(q) - b(q) < 0$ , and b(q) < 0, and  $\Delta D(q) \equiv D_2(q) - D_1(q) \ge 0$ , and also that  $D_1(q)$ ,  $\Delta D(q)$ , and d(q) satisfy

$$\frac{2D_1(q) + \mu d(q)}{\Delta D(q) + (\mu + 1)d(q)} > \frac{b(q)}{k^2 - a(q) - b(q)}.$$
(7)

Then it follows that  $\Delta(k,q) \leq 0$ .

- For dissipation term configurations, choose one of
  - For  $\mu = 0$  and  $\lambda = 1$ : standard Ekman at  $p_s = p_2$ .
  - For  $\mu = -1/3$  and  $\lambda = 3/2$ : extrapolated Ekman at  $p_s = 1$ Atm.
- and also one of
  - For  $\Delta \nu > 0$ : differential small-scale diffusion
  - For  $\Delta \nu = 0$ : no differential small-scale diffusion

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#### The two-layer model. VI

For μ = 0 and λ = 1: standard Ekman at p<sub>s</sub> = p<sub>2</sub>; and Δν = 0: no differential diffusion

$$\frac{\nu_E}{4\nu q^{2p}} \leq \frac{q^2 - k^2}{k_R^2} \Longrightarrow \Delta(k,q) \leq 0.$$

For μ = −1/3 and λ = 3/2: extrapolated Ekman at p<sub>s</sub> = 1Atm; and Δν > 0: with differential diffusion

$$0 < \frac{\Delta\nu q^{2p} + \nu_E}{4\nu q^{2p} - \nu_E} < \frac{q^2 - k^2}{k_R^2} \Longrightarrow \Delta(k, q) \le 0.$$
(8)

- ► Note that the hypothesis requires that v<sub>E</sub> < 4vq<sup>2p</sup> (thank extrapolated Ekman)
- Increasing either ν<sub>E</sub> or Δν indicates a tendency towards violating the flux inequality.
- For Δν > 0, the LHS of hypothesis will approach Δν/(4ν) and remain bounded for large wavenumbers q (thank differential diffusion)

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#### The two-layer model. VII

For μ = 0 and λ = 1: standard Ekman at p<sub>s</sub> = p<sub>2</sub>; and Δν > 0: with differential diffusion

$$\frac{\Delta\nu q^{2p} + \nu_E}{4\nu q^{2p}} < \frac{q^2 - k^2}{k_R^2} \Longrightarrow \Delta(k, q) \le 0, \tag{9}$$

- Hyperbolic blow-up is no longer possible.
- For Δν > 0, the LHS of hypothesis will still approach Δν/(4ν) and remain bounded for large wavenumbers q
- ► For  $\mu = -1/3$  and  $\lambda = 3/2$ : **extrapolated** Ekman at  $p_s = 1$ Atm; and  $\Delta \nu = 0$ : **no** differential diffusion

$$\frac{\nu_E}{4\nu q^{2p}} < \frac{q^2 - k^2}{k_R^2 + (q^2 - k^2)} \Longrightarrow \Delta(k, q) \le 0.$$
<sup>(10)</sup>

- ▶ LHS vanishes with increasing *q* but RHS remains bounded.
- Condition is still tighter.

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## The two-layer model. VIII

- Sufficient conditions can be derived in terms of the streamfunction spectra U<sub>1</sub>(q) = ⟨ψ<sub>1</sub>, ψ<sub>1</sub>⟩<sub>k</sub>, U<sub>2</sub>(q) = ⟨ψ<sub>2</sub>, ψ<sub>2</sub>⟩<sub>k</sub>, and C<sub>12</sub>(q) = ⟨ψ<sub>1</sub>, ψ<sub>2</sub>⟩<sub>k</sub>
- ► Arithmetic-geometric mean inequality:  $2|C_{12}(q)| \le U_1(q) + U_2(q)$
- ▶ PROPOSITION 3: Assume streamfunction dissipation with  $\Delta \nu = 0$ , and standard Ekman (i.e.  $\mu = 0$  and  $\lambda = 1$ ) with d(k) > 0 and  $k^2 a(q) b(q) < 0$  and b(q) < 0 and  $C_{12}(q) \le U_2(q)$ . Then, it follows that  $\Delta(k, q) \le 0$ .
- ▶ PROPOSITION 4: Assume that b(q) < 0 and k<sup>2</sup> − a(q) − b(q) < 0. Assume also streamfunction dissipation with both differential small-scale dissipation and extrapolated Ekman dissipation with −1 < µ < 0.</p>
  - 1. If  $C_{12}(q) \le 0$ , then  $\Delta(k,q) \le 0$ .
  - 2. If  $C_{12}(q) \le \min\{U_1(q), U_2(q)\}$  and  $U_1(q) + \mu U_2(q) \ge 0$ , then  $\Delta(k,q) \le 0$ .
- ▶ PROPOSITION 5: Assume that k<sup>2</sup> − a(q) − b(q) < 0 and b(q) < 0. We also assume streamfunction dissipation with extrapolated Ekman dissipation with −1 < μ < 0 and symmetric small-scale =</li>

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## Conclusions

- Under symmetric dissipation, the flux inequality is satisfied unconditionally for multi-layer QG models
- Under asymmetric dissipation, the flux inequality is satisfied only when the asymmetry satisfies restrictions.
- The restrictions given are sufficient but not necessary.

# Thank you!

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